

A progress report on the estimation of geoacoustic parameters from measured time series taken during the East China Sea component of ASIAEX

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OUTLINE

- I. Modeling of forward transmission data
- II. Comparisons to analyses of previous experiments
- III. Summary and future work



Issues and Objectives

Extract information on physics of low-frequency scattering from seabed

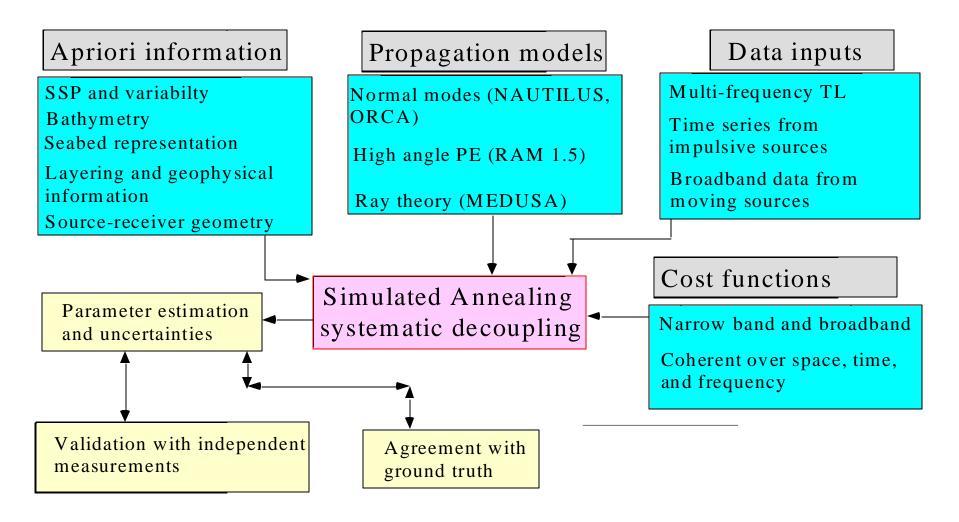
Use acoustic forward transmission data to establish average geoacoustic structure

Employ this information in the scattering equations to invert for the scattering potential of the seabed using reverberation data.

What features in the forward data allow for extraction of sound speed and attenuation structure?

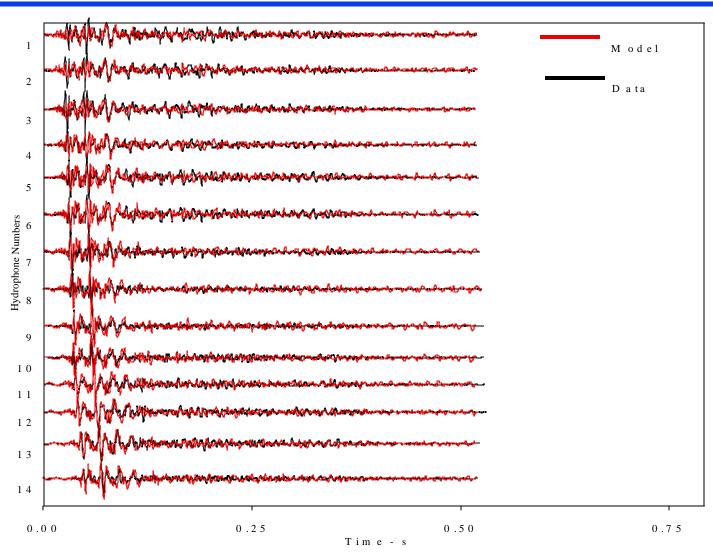


ARL Analysis and Inversion Methodology



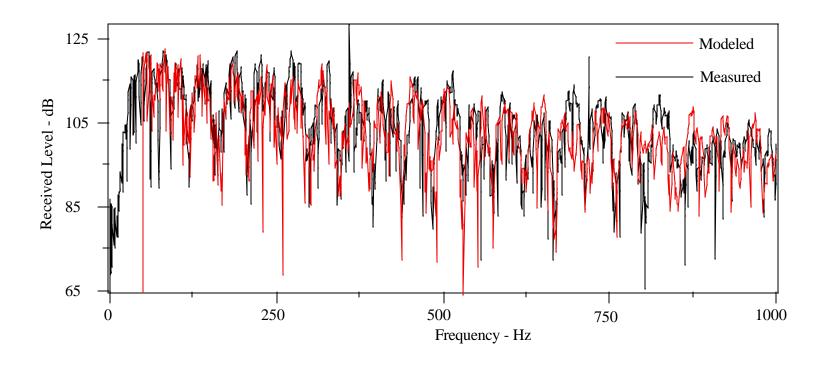






Comparison of measured and modeled time series data on the APL/URI VLA. Range = 16.94 km, source depth = 55 m





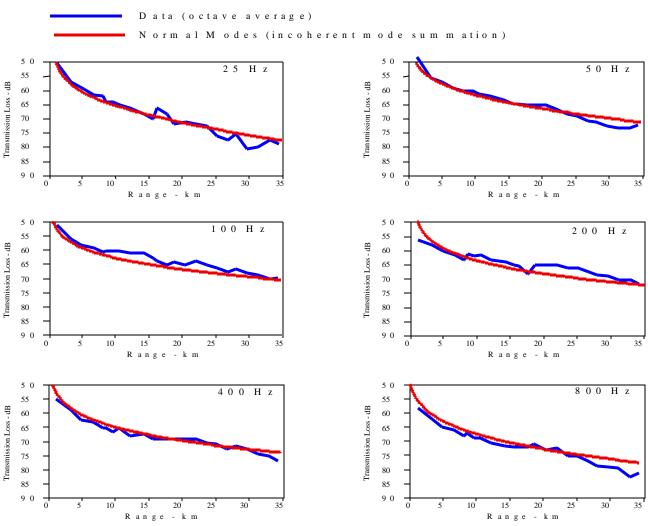
Comparison of measured and modeled magnitude of spectra for a single phone on VLA in East China Sea measurements. Source-receiver range = 16.94 km. Source depth 55 m. Source spectra from Chapman-Wakely model convolved with modeled frequency response in 50-350 Hz band. Modeled frequency response made with ORCA normal mode model.



Depth - m	m Compressional Speed - m/sDensity - g/cc		Attenuation db/m-kHz Frequency Exponent		
0.0	1640.0	1.76	0.45	1.8	Sand
7.6	1664.9	1.76	0.45	1.8	
7.6	1839.8	1.9	0.27	1.0	Sand-Gravel
95.5	1918.1	1.9	0.27	1.0	
95.5	3000.0	2.5	0.02	1.0	Basement

Current Estimated Geoacoustic Profile in ECS basin from data analysis





Comparison of modeled TL using the ECS geoacoustic profile to me TL taken at 29 05 N 126 43 E.





Observations

Estimated sound speed structure of upper 10 m of sediment describes relative arival times of main multipath arrivals and general time spread characteristics as a function of range. Arrivals demonstrate critical angle effect at the water sediment interface.

A top sediment sound speed of 1640 m/s with a high sound speed depth gradient appears to be the key structure of the geoacoustic profile. Analysis, however, may have slightly overestimated sound speed because of incorrect water depth.

Very difficult to establish precise sediment thickness from time series data.

ARL Comparisons of geoacoustic profiles in ECS region

East China Sea basinYellow Sea basinSouthern Strait of Kore*

Depth - m	Compressional Speed - m/s	Density - g/cc	Attenuation db/m-kHz Frequency Exponen	
0.0 0.0 0.0	1640.0 1644.7 1629.0	1.76 1.72 1.88	0.45 0.89 0.28	1.8 2.0 1.8
7.62 5.96 10.0	1664.9 1650.4 1644.0	1.76 1.72 1.88	0.45 0.94 0.30	1.8 2.0 1.8
7.62 5.96 10.0	1839.8 1725.3 1760.0	1.9 1.8 1.95	0.27 0.47 0.15	1.0 1.3 1.8
95.48131.9 132.0	1918.1 1848.9 1761.0	1.9 1.8 1.95	0.27 1.0 0.15	1.0 1.3 1.8
95.48131.1 132.0	3000.0 3000.0 2200.0	2.5 2.5 2.2	0.02 0.02 0.02	1.0 1.0 1.0

Geoacoustic Models for ECS and YS basins and SOK
Obtained From Analysis and Inversion of TL and
Time Series Data.

* Rozenfeld et al IEEE JOE

Knobles et al

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Summary

An approximate geoacoustic description has been established for ECS Consistent with previous geophysical measurements.

Key characteristics of both forward time series and transmission loss measurements are reproduced by range-independent

Geoacoustic model used in analysis of scattering data.

Uniquess of attenuation and layering structure in seabed difficult to es

Elastic description may be required for consistent picture of seabed acoustics.



Modeling of ASIAEX Reverberation

Research in progress



Sources of Reverberation



$$z_{\rm int} = H + \boldsymbol{z}(r, \boldsymbol{q})$$

Volume Scattering

$$c_{sed} = c_s(z)[1+\boldsymbol{m}(r,\boldsymbol{q},z)]$$

$$\boldsymbol{r}_{sed} = \boldsymbol{r}_{s}(z)[1+\boldsymbol{h}(r,\boldsymbol{q},z)]$$

Model Assumptions

- Density fluctuations proportional to sound speed fluctuations: $h(\mathbf{r}) = 2gm(\mathbf{r})$
 - Sound speed fluctuations small: $|\mathbf{m}(\mathbf{r})| << 1$
- Water sediment interface nearly flat, slightly rough: $|k_0 \mathbf{z}(\mathbf{r})| << 1 \quad |\nabla_{\parallel} \mathbf{z}(\mathbf{r})| << 1$

• Perturbation approximation: keep only first order terms in above small quantities

Equation for Scattered Field

$$p_{scat}(\mathbf{r}) = -\frac{1}{4p} \begin{cases} \int_{V} (2\mathbf{g} + 2)\mathbf{m}(\mathbf{r}')k_{0}^{2}(z')[G_{0}(\mathbf{r}; \mathbf{r}')G(\mathbf{r}'; \mathbf{r}_{0})]dV' \\ -\int_{V} (2\mathbf{g})\mathbf{m}(\mathbf{r}')[\nabla'G_{0}(\mathbf{r}; \mathbf{r}') \cdot \nabla'G(\mathbf{r}'; \mathbf{r}_{0})]dV' \\ +\int_{S} (2\mathbf{g})\mathbf{m}(\mathbf{r}'_{-}) \left[G_{0}(\mathbf{r}; \mathbf{r}'_{+}) \frac{\partial G}{\partial z'}(\mathbf{r}'_{+}; \mathbf{r}_{0})\right]dA' \\ +\int_{S} \mathbf{z} (\mathbf{r}') \left(1 - \frac{\mathbf{r}_{w}}{\mathbf{r}_{s}}\right) \left[\nabla'_{\parallel} G_{0}(\mathbf{r}; \mathbf{r}'_{+}) \cdot \nabla'_{\parallel} G(\mathbf{r}'_{+}; \mathbf{r}_{0})\right]dA' \\ -\int_{S} \mathbf{z} (\mathbf{r}')\mathbf{w}^{2} \left(\frac{1}{c_{w}^{2}} - \frac{\mathbf{r}_{w}}{\mathbf{r}_{s}c_{s}^{2}}\right) G_{0}(\mathbf{r}; \mathbf{r}'_{+}) G(\mathbf{r}'_{+}; \mathbf{r}_{0})]dA' \end{cases}$$

Fluctuation Fields

Described by radially isotropic, power-law, power spectra:

$$W(k_r) = \frac{4pns^2 l_r^2}{(1 + l_r^2 k_r^2)^{n+1}}$$

$$W(k_r) = \frac{8p^{3/2}s^2 l_r^2 l_z \Gamma(n + 3/2)}{\Gamma(n)(1 + l_r^2 k_r^2 + l_z^2 k_z^2)^{n+1}}$$

Simulation Details

- Single-scattering: Use Born Approximation (i.e. $G(\mathbf{r}_1; \mathbf{r}_2) = G_0(\mathbf{r}_1; \mathbf{r}_2)$)
- Monostatic scattering geometry
- Compute unperturbed Green's function by normal mode decomposition
- Narrow (30 Hz) band processing
- Fit ensemble-averaged simulated time series to band-filtered time series via power spectrum parameters



References

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